

# Sea level and ocean/land/ice observations and models

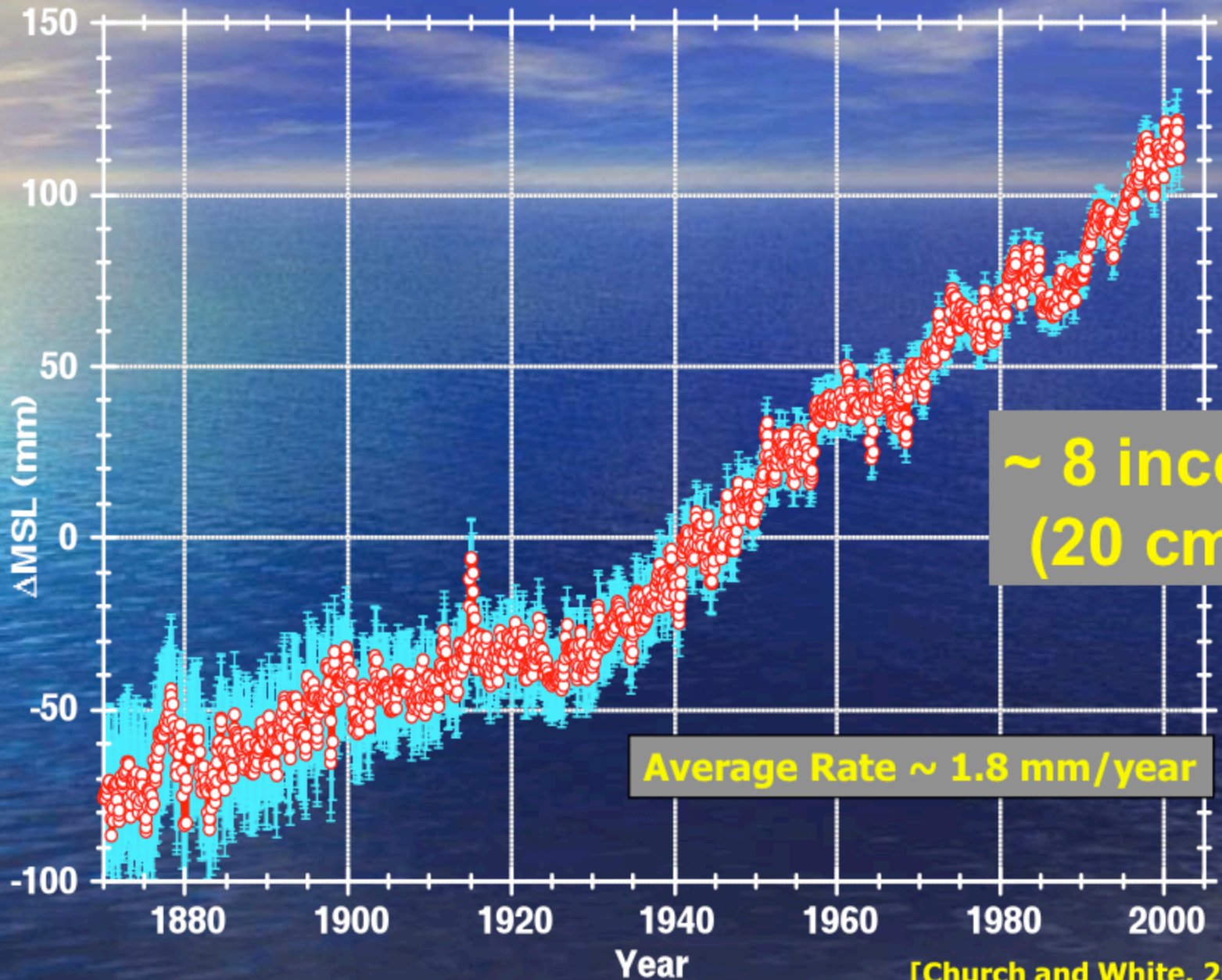
Science discussions with Mike Freilich

Jet Propulsion Laboratory

December 10, 2010

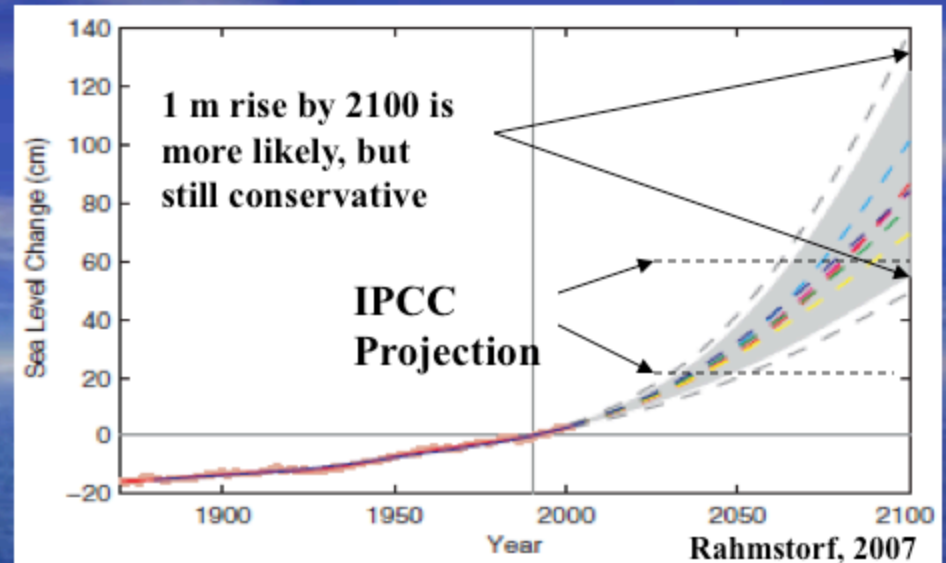
- Sea level and altimetry (Lee-Lueng Fu)
- Land/ice observations and models (Erik Ivins)
- ISSM: ice sheet data assimilation (Eric Larour)
- ECCO2: ocean and sea ice data assimilation (Dimitris Menemenlis)
- Sea ice observations and models (Ron Kwok)
- Ice sheet/ocean interactions (Eric Rignot)

# NASA Owns the Sea Level Problem





# Prediction of sea level and its geographic patterns is a grand challenge for the coming decades



## Combining Data to Study Sea Level Change

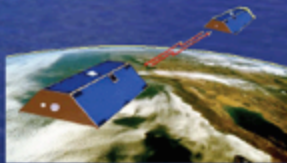
Addition of Heat



Argo

+

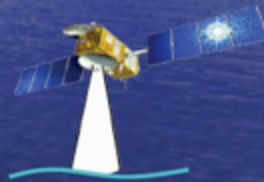
Addition of Freshwater



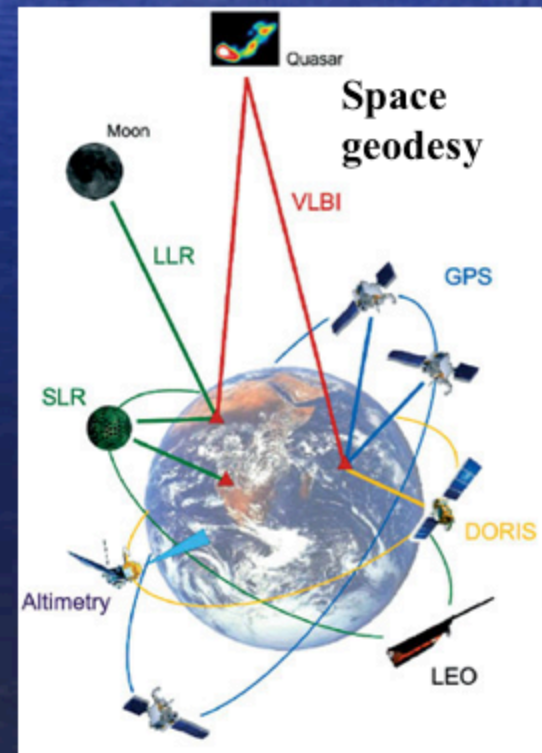
GRACE

=  
(roughly)

Total Sea Level Rise



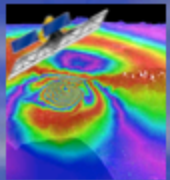
Jason



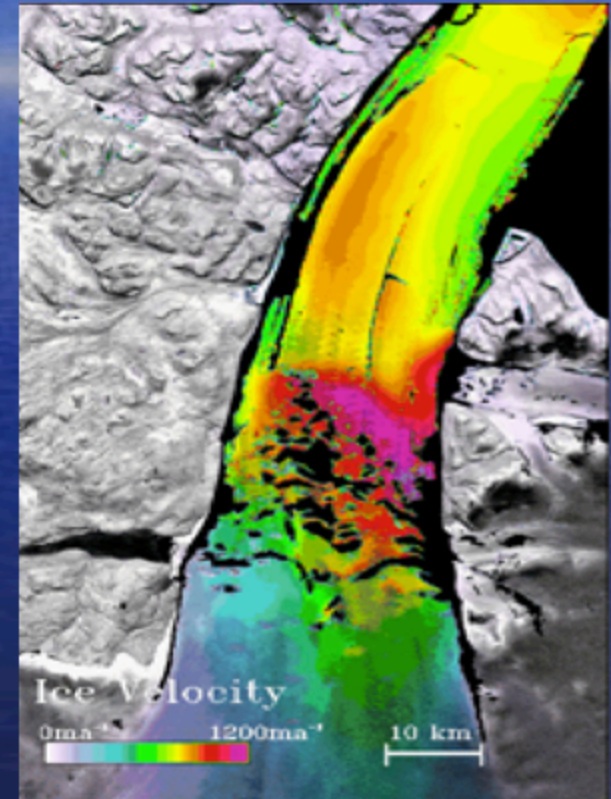


# Melting Ice Sheet is the Wild Card...

- Repeat-pass radar interferometry allows us to measure how quickly ice sheets are moving from space

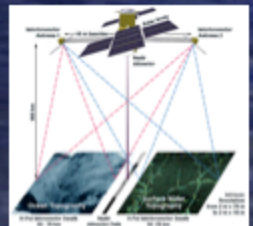


DESDynI



## Ice shelves holding up the ice sheets

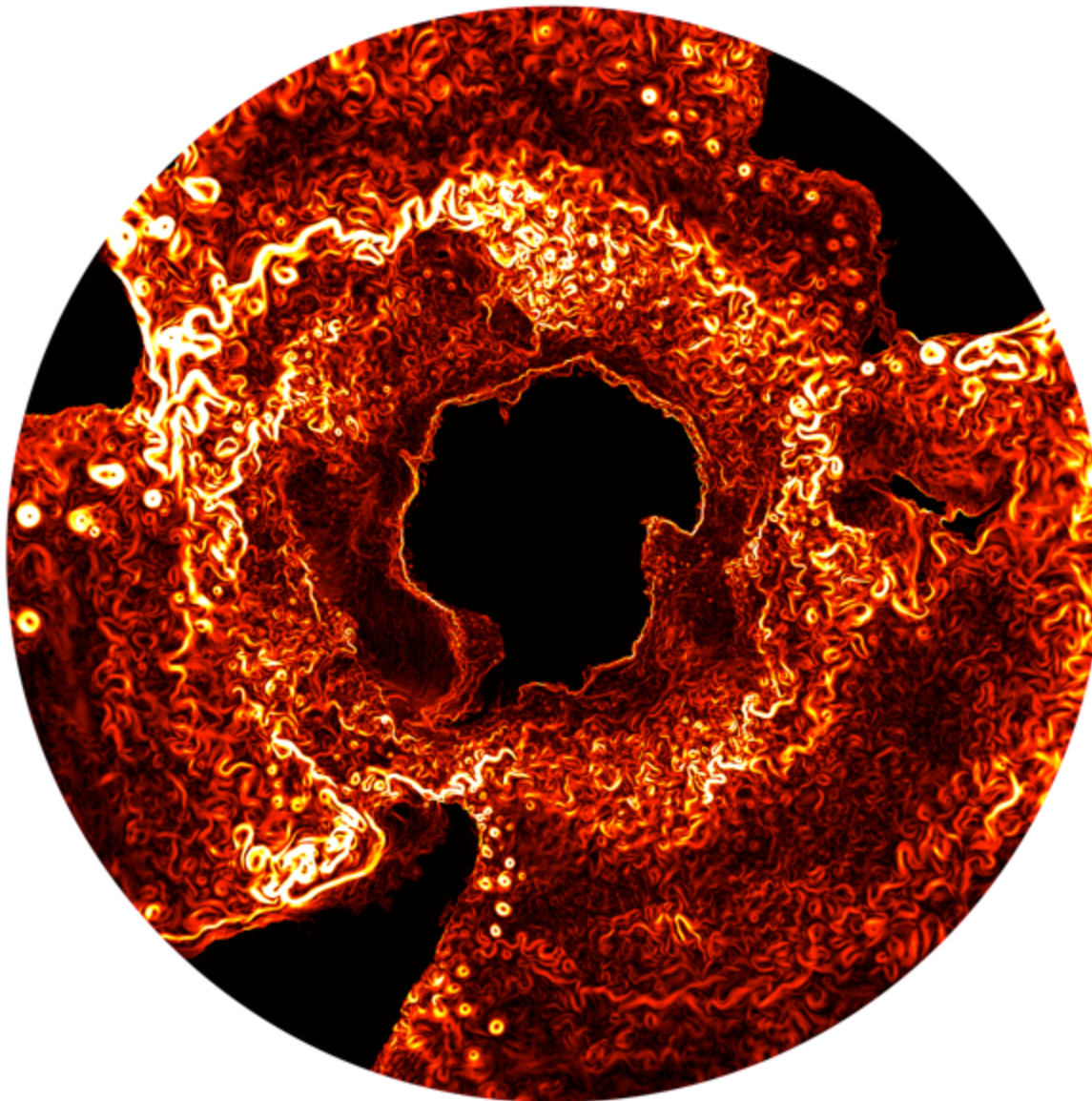
- Wide-swath altimetry allows us to measure the ocean currents transporting heat to melt the ice shelves



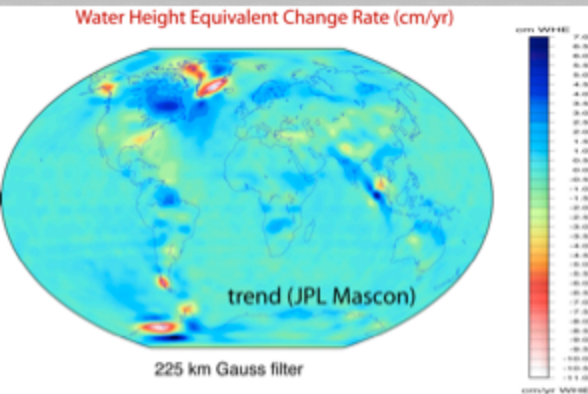
SWOT



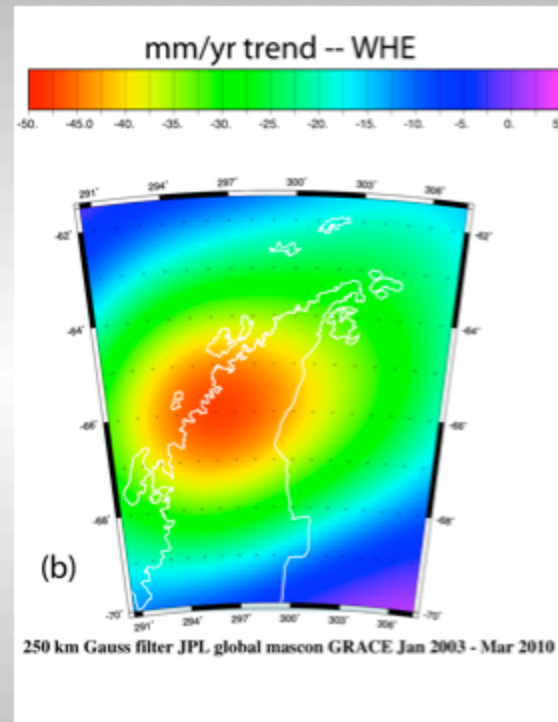
# JPL Ocean-Ice Modeling Effort



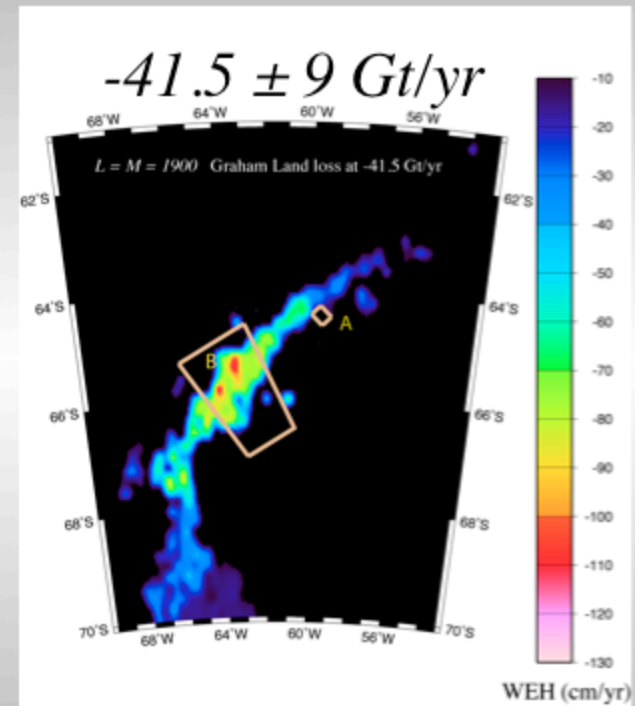
# 8-yr GRACE ice loss record



*GRACE trends over 8 Years: Deduction of Cryospheric to Ocean Mass Transport: Roughly large enough to explain ongoing sea level rise*



*JPL Global Mascon solution  
At the Antarctic Peninsula*



*Model ice and solid  
Earth Isostatic response  
(GPS data) & apply smoothing -  
compare to altimetry*

*Antarctica -> 220-246 Gt/yr  
(Chen et al. 2009)*

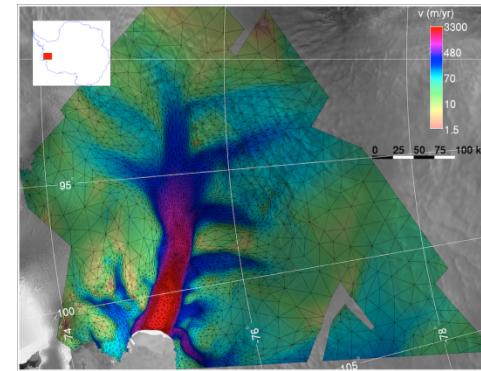
*Antarctica ->  $100 \pm 67 \text{ Gt/yr}$   
(GIA corrected: Riva et al., 2009)*

*Greenland -> 137-159 accelerated to 267-28 Gt/yr  
(Velicogna, 2009)*

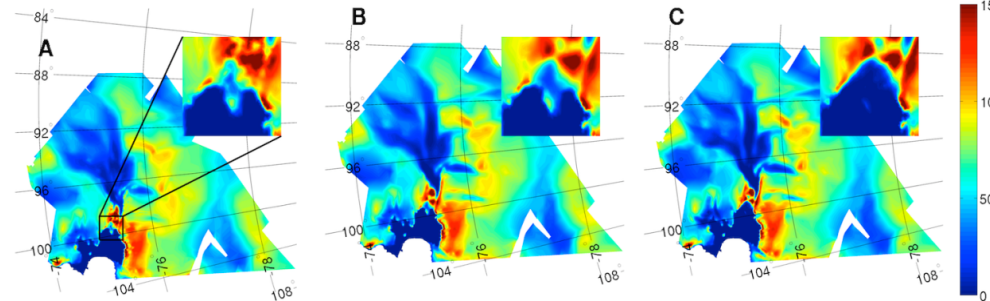


# ISSM: Ice Sheet System Model.

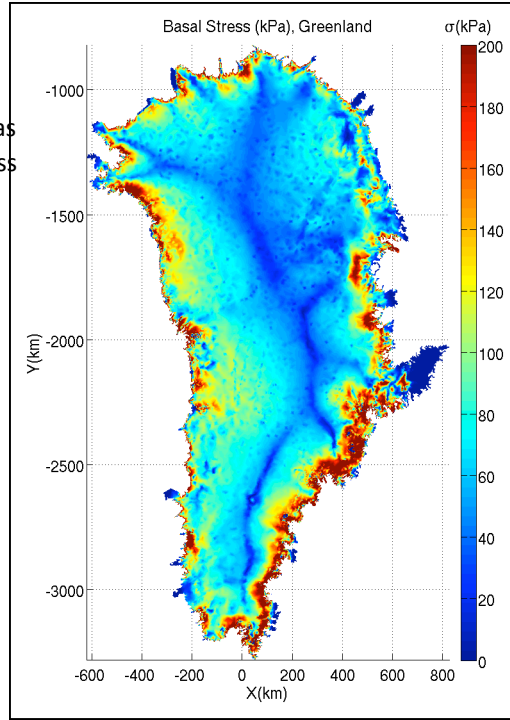
- finite element model, anisotropic meshing -> refine model where physics are warranted.
- large scale capabilities: high resolution (1km horizontal, 10 layers vertical).
- multi-model: 2D, 3D higher-order and 3D Full-Stokes
- data assimilation: optimize unknown parameters ( basal stress and ice rigidity) using the adjoint model (inverse control method) and surface velocity data from InSAR.



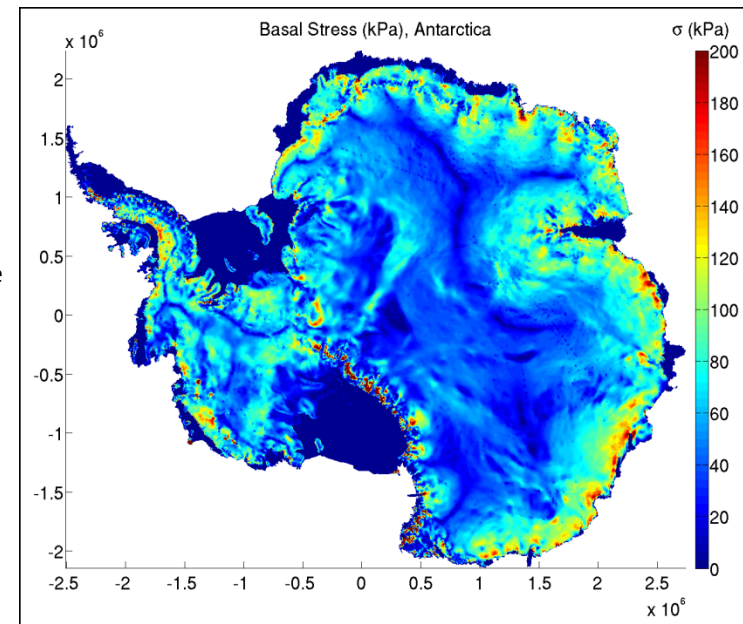
Pine Island Glacier: assimilation of basal stress using 2D (frame A), 3D Blatter/Pattyn (frame B) and 3D Full-Stokes (frame C) models. Basal stress near the grounding line is extremely model dependent -> need for Full-Stokes locally. Morlighem et al 2010.



Large scale inversion of basal stress over the Greenland Ice Sheet. A Full-Stokes model was used to inverse the basal stress using control methods and InSAR surface velocities from Joughin 2010.

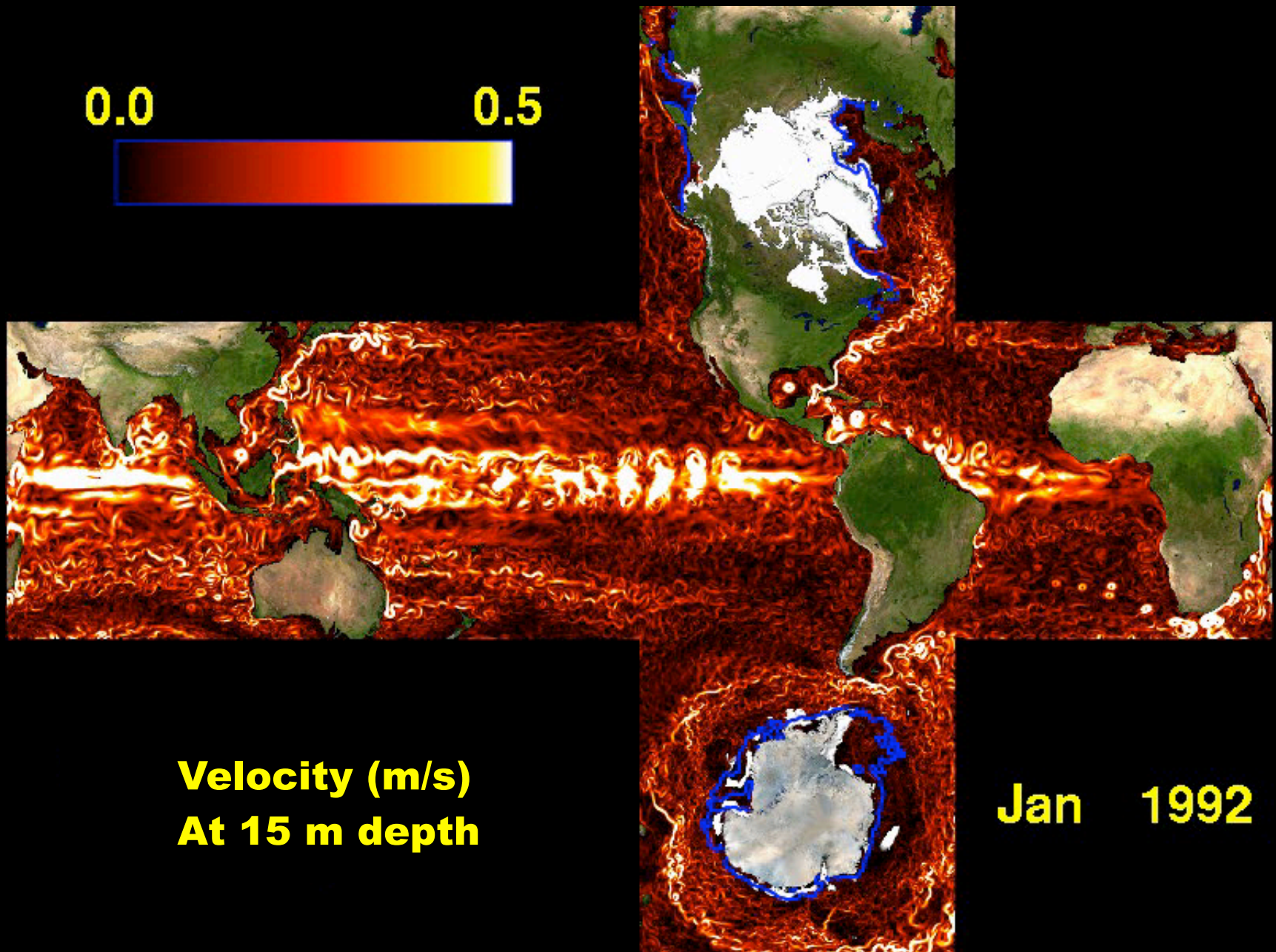


Large scale inversion of basal stress over the Antarctica Ice Sheet. A 3F Blatter/Pattyn model was used to invert the basal stress using control methods and InSAR surface velocities (Rignot, unpublished). Larour, GRL 2010 in revision.



# ECCO2

Eddyding global-ocean and sea-ice data synthesis for improved estimates and models of ocean carbon cycle, understanding recent evolution of polar oceans, monitoring time-evolving term balances within and between different components of the Earth system, and many more science applications.

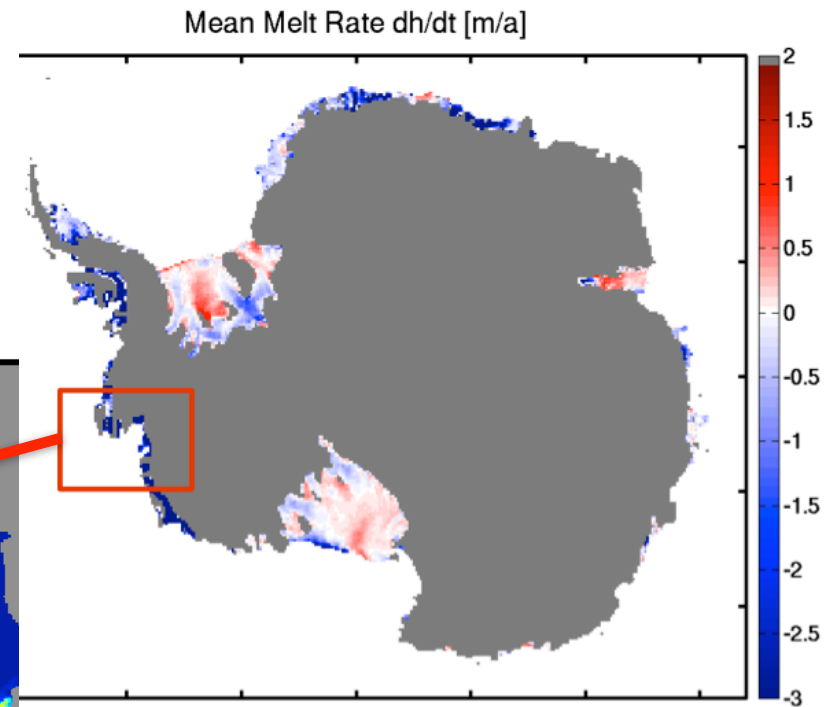
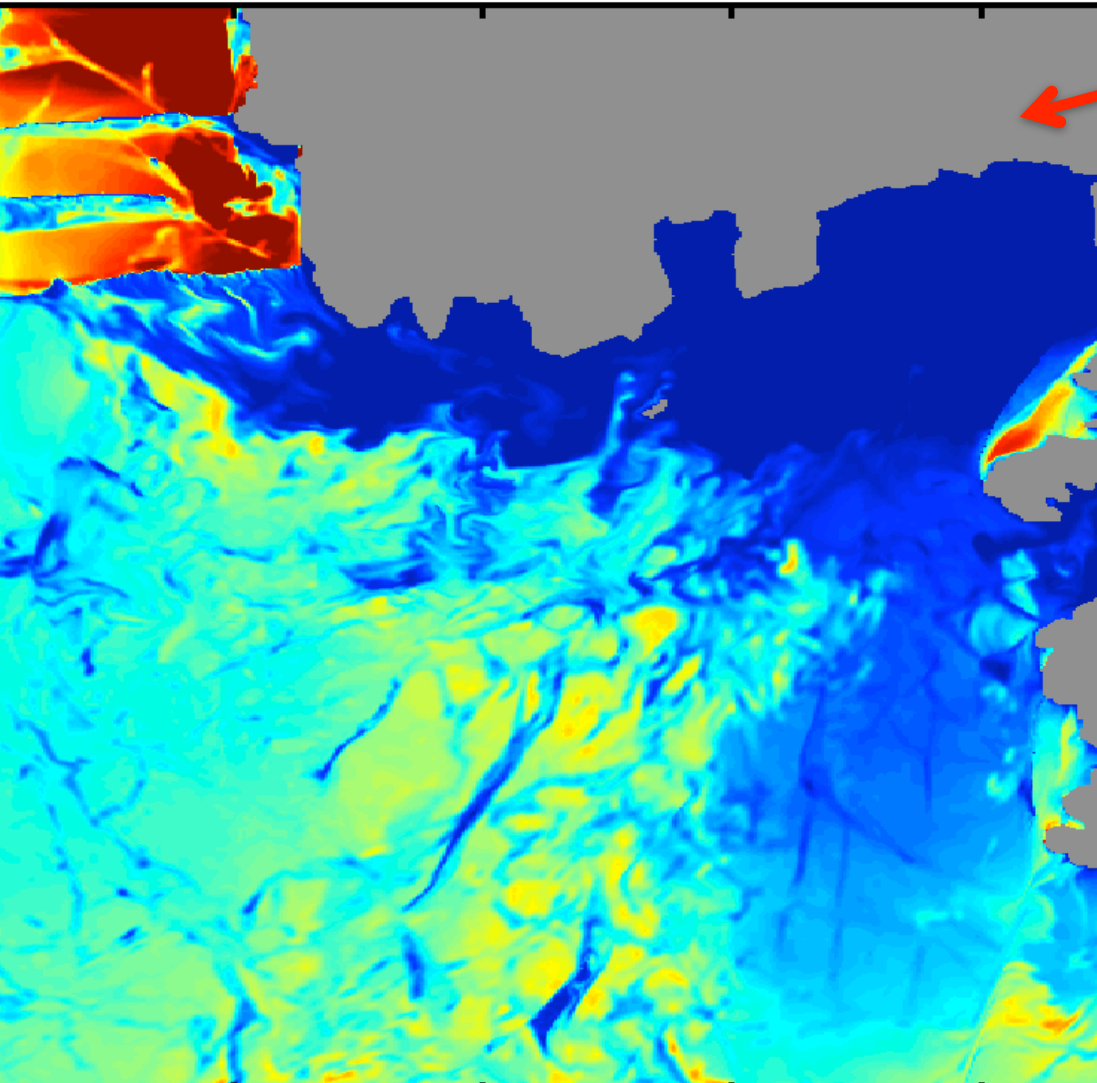




# Ice shelf cavities in ECCO2

(Michael Schodlok, JPL/JIFRESSE)

Sea Ice Thickness: 01-Jan-1980



**ECCO2 basal melt estimate**  
(59 mSv) is double previous  
(BRIOS) estimate, more  
consistent with mass loss  
derived from ICESat/GLAS.

**Sea ice in the coupled system**  
is at the heart of the “ice  
pump” mechanism governing  
sub-ice shelf cavity circulation  
and an essential ingredient in  
the production of very dense  
High Salinity Shelf Water, which  
is the source of Antarctic  
Bottom Water.

# Observations of sea ice kinematics

(Ron Kwok and Gunnar Spreen, JPL)

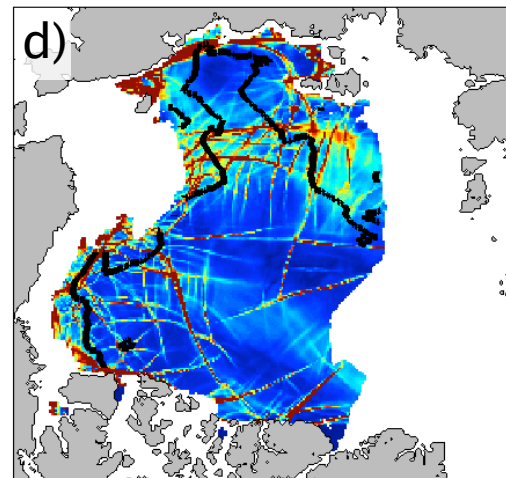
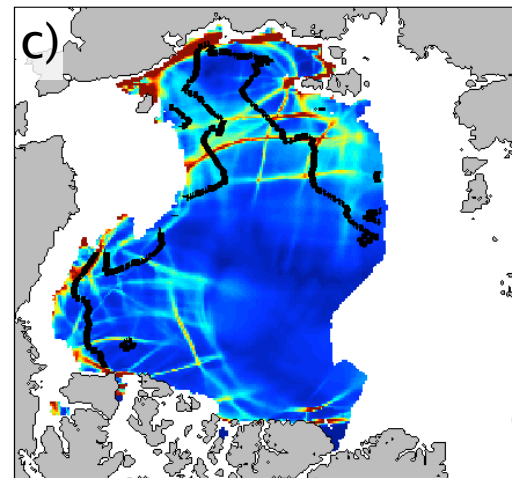
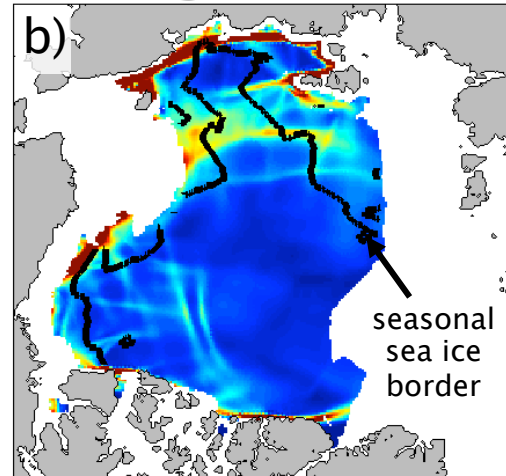
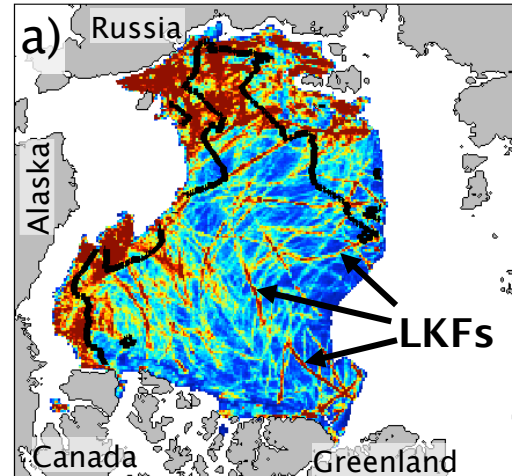


# New sea ice constitutive model

(Deborah Sulsky, UNM and Ron Kwok, JPL)

RGPS

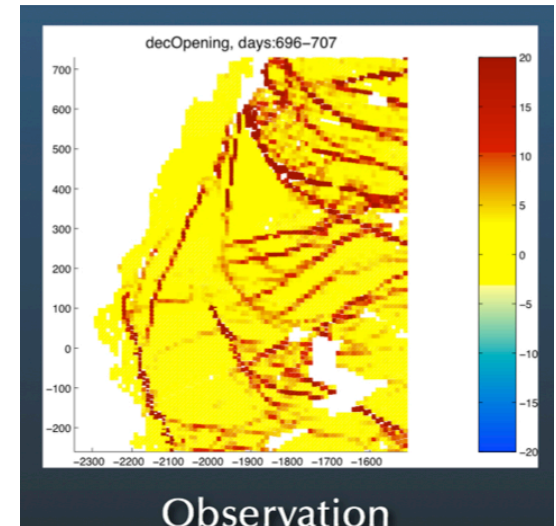
MITgcm 18 km



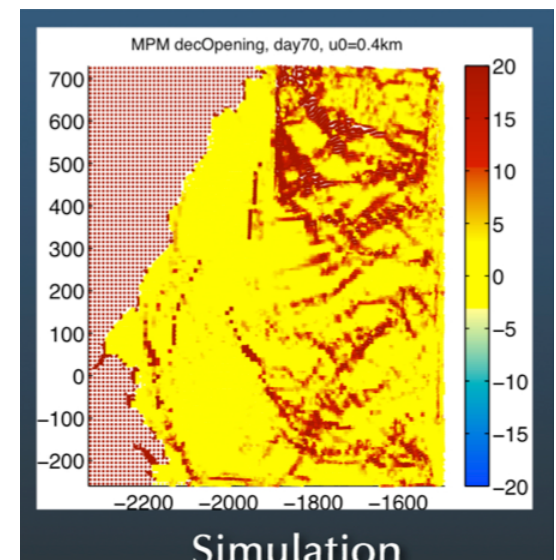
MITgcm 9 km

MITgcm 4 km

Nov. 1997 sea ice deformation:  $D = \sqrt{\text{div}^2 + \text{shear}^2}$



Observation

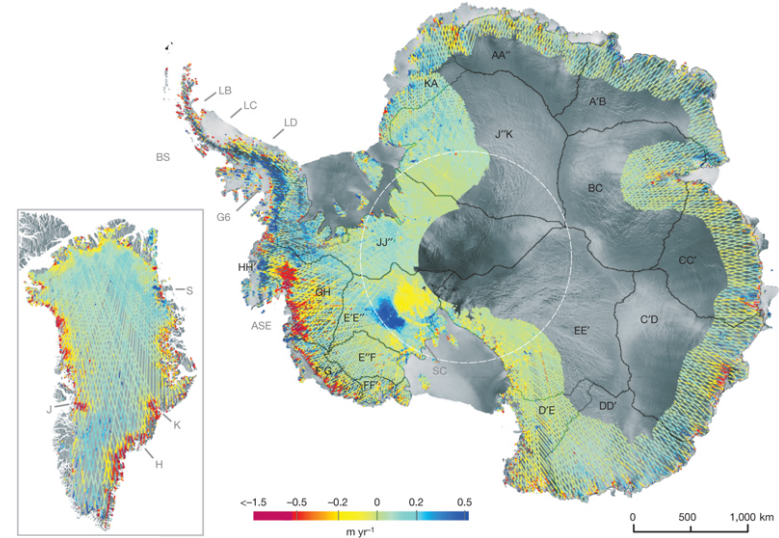
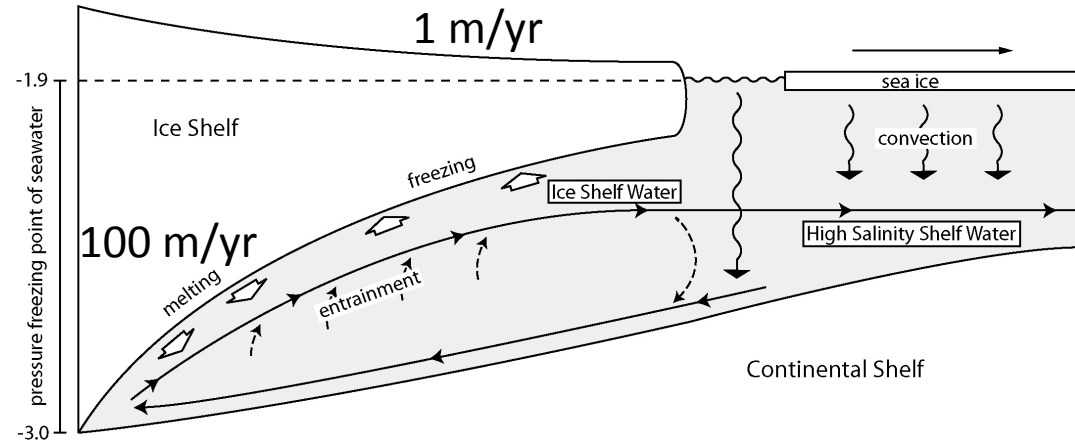


Simulation

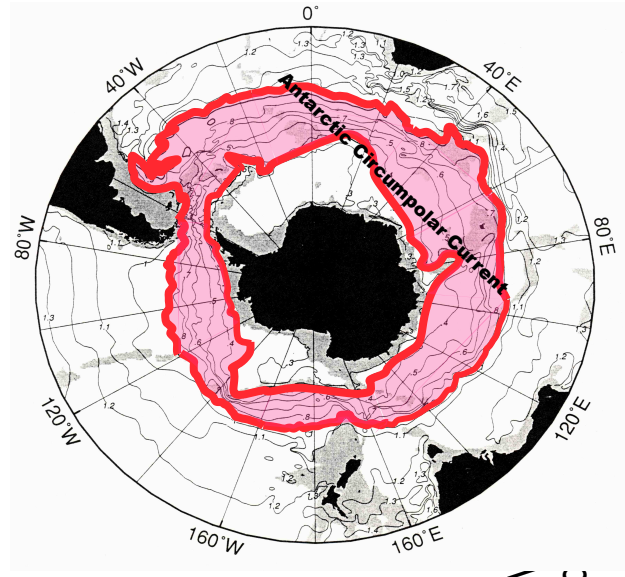
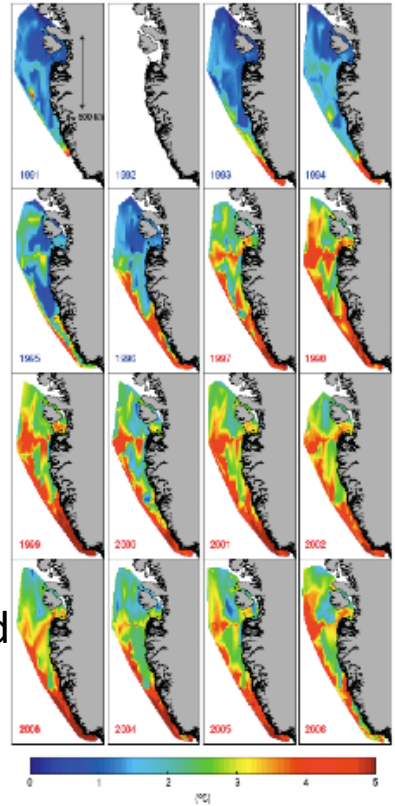
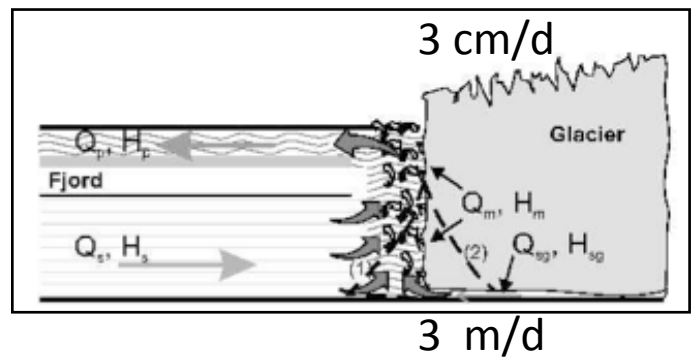


# Ice-sheet/Ocean Interactions

Rapid ice sheet changes at low elevation



Sub-aqueous melt rates > 100 surface melt rates



- Surface mass balance steady in Antarctica but declining 100% by decade in Greenland due to enhanced melt.
- Yet surface melt alone cannot explain rapid changes in glacier dynamics.
- Why? Ice sheets melt from the bottom

Significant oceanic changes but impact on land ice is poorly documented

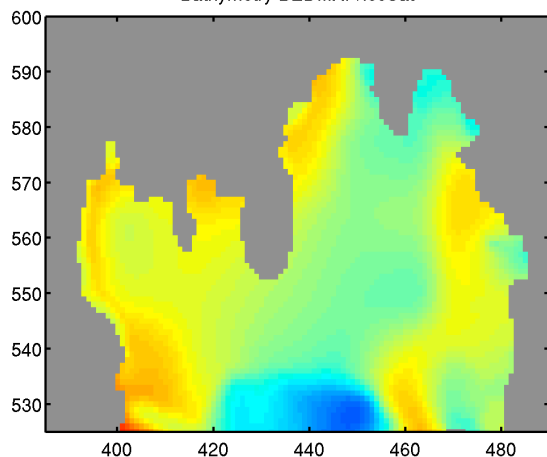


NASA plays a critical role:

- IceBridge: dhdt, H, bathymetry
  - ECCO2 model/data assimilation
  - ISSM model/data assimilation
- high-res, continental view.

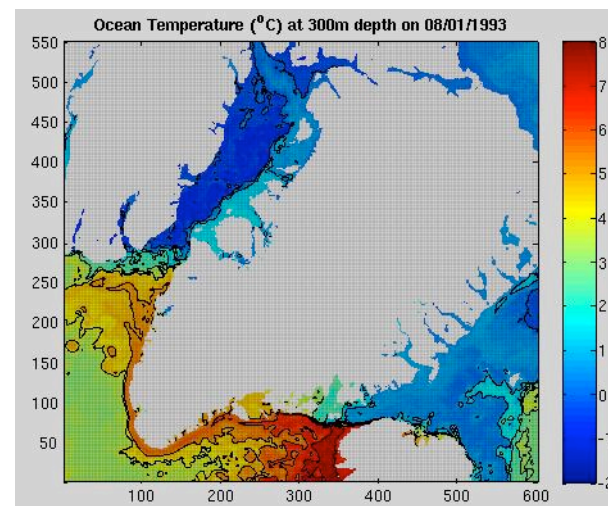
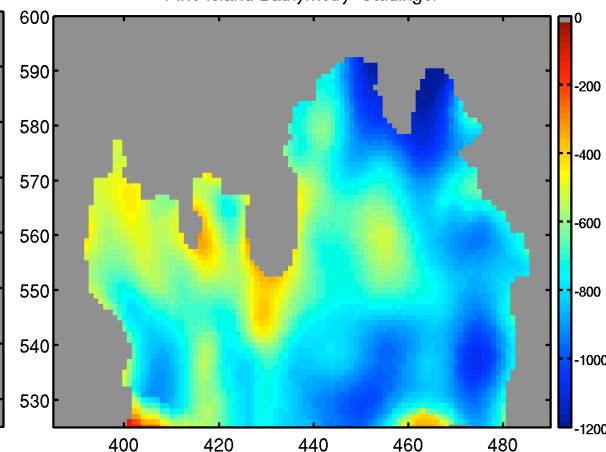
Unrealistically low melt

Bathymetry BEDMAP/iceSat



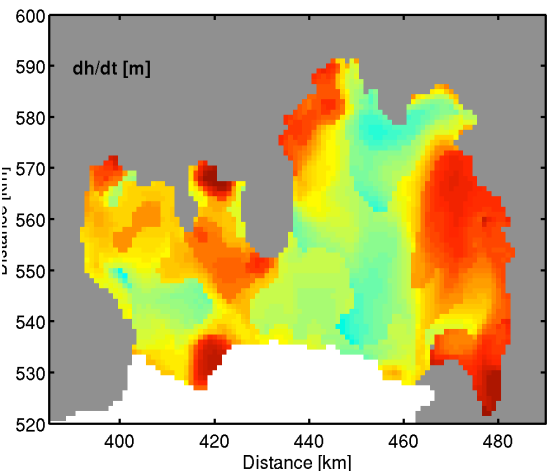
Consistent with satellite estimates

Pine Island Bathymetry Studinger



Oceanic change migrates northwest

Mean Melt Rate Pine Island 1980



Mean Melt Rate Pine Island 1980

